Simulating forest productivity and surface-atmosphere carbon exchange in the BOREAS study region

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Summary A process-based, general ecosystem model (BI-OME-BGC) was used to simulate daily gross primary production, maintenance and heterotrophic respiration, net primary production and net ecosystem carbon exchange of boreal aspen, jack pine and black spruce stands. Model simulations of daily net carbon exchange of the ecosystem (NEE) explained 51.7% (SE = 1.32 g C m⁻² day⁻¹) of the variance in daily NEE derived from stand eddy flux measurements of CO2 during 1994. Differences between measured and simulated results were attributed to several factors including difficulties associated with measuring nighttime CO2 fluxes and model assumptions of site homogeneity. However, comparisons between simulations and field data improved markedly at coarser timescales. Model simulations explained 66.1% (SE = 0.97 g C m⁻² day⁻¹) of the variance in measured NEE when 5-day means of daily results were compared. Annual simulations of aboveground net primary production ranged from 0.6-2.4 Mg C ha⁻¹ year⁻¹ and were concurrent with results derived from tree increment core measurements and allometric equations.

Model simulations showed that all of the sites were net sinks (0.1-4.1 Mg C ha⁻¹ year⁻¹) of atmospheric carbon for 1994. Older conifer stands showed narrow margins between uptake of carbon by net photosynthesis and carbon release through respiration. Younger stands were more productive than older stands, primarily because of lower maintenance respiration costs. However, all sites appeared to be less productive than temperate forests. Productivity simulations were strongly linked to stand morphology and site conditions. Old jack pine and aspen stands showed decreased productivity in response to simulated low soil water contents near the end of the 1994 growing season. Compared with the aspen stand, the jack pine stand appeared better adapted to conserve soil water through lower daily evapotranspiration losses but also exhibited a narrower margin between daily net photosynthesis and respiration. Stands subjected to water stress during the growing season may exist on the edge between being annual sources or sinks for atmospheric carbon.

Keywords: BIOME-BGC, boreal forest, carbon balance, ecosystem modeling, net primary production.

Introduction

The boreal forest covers a broad circumpolar band across the Eurasian and North American continents and represents approximately 11% of the Earth's total land area (Bonan and Shugart 1989). Global climate simulation studies show that the boreal forest region should experience significant warming and drying in response to increases in atmospheric CO₂ concentration and other greenhouse gases (Manabe and Stouffer 1980, Houghton et al. 1990, Monserud et al. 1993). This response is of major concern because the boreal forest region is thought to contain 16–24% of the world's soil carbon (Gates 1993). Significant warming in this region may result in fairly rapid, large-scale displacement and redistribution of boreal forests, enhanced release of CO₂ to the atmosphere and intensification of global warming (Emanuel et al. 1985, Pastor and Post 1988, Neilson and Marks 1994).

A critical goal of the BOREAS project is to integrate point measurements across multiple spatial and temporal scales using process level models of the boreal forest water, energy and biogeochemical cycles (BOREAS Science Steering Committee 1990). BIOME–BGC is a general ecosystem process model designed to simulate daily biogeochemical and hydrologic processes from stand to global scales (Running and Hunt 1993). Our goal was to utilize BIOME–BGC logic to synthesize field measurements and establish linkages between measured data at various spatial and temporal scales. Specifically, we used BIOME–BGC to simulate daily and annual carbon budgets of several boreal forest stands and compare net carbon flux simulations with results derived from tower eddy flux and biomass measurement data. We also investigated climate and stand characteristics that control simulated carbon fluxes.

Experimental design

BIOME--BGC

BIOME–BGC simulates biogeochemical and hydrologic processes across multiple biomes based on the logic that differences in process rates among biomes are primarily a function of climate and general life-form characteristics. The carbon balance portion of BIOME–BGC utilizes daily meteorological

data in conjunction with general stand and soil information to predict net photosynthesis, growth, maintenance and heterotrophic respiration at a daily time-step. BIOME-BGC is general in the sense that the surface is represented by singular, homogeneous canopy and soil layers. The model is similar in scope and logic to the FOREST-BGC model described by Running and Coughlan (1988), except that site and vegetation components have been determined and tested for multiple biomes, including coniferous and deciduous forests, grassland, shrub-land and alpine cover types (Running and Hunt 1993, Cairns and Mananson 1996, Kremer and Running 1996). The boreal forest represents a unique environment characterized by flat terrain and a short growing season. Measurements from other BOREAS investigators have shown large sensible heat fluxes during the growing season resulting from low solar albedos and low evapotranspiration rates for coniferous stands (Sellers et al. 1995). This investigation constitutes our initial effort to test and improve BIOME-BGC logic at the stand level in order to improve model representation of the boreal forest biome in regional applications. Because BIOME-BGC logic has been described in detail by Running and Coughlan (1988) and Running and Hunt (1993), we present here only a summary of the important components of BIOME-BGC relating to the prediction of daily carbon allocation and exchange.

Gross primary production (GPP) represents the total gain of carbon to the system by net photosynthesis and is defined as the daily sum of gross photosynthesis and daily foliar respiration $(R_{\rm dl})$. The current representation of photosynthesis differs significantly from previously published descriptions of the BGC family of models. The current implementation of photosynthetic biochemistry is closely related to the original BI-OME-BGC logic in that it is based on the Farquhar biochemical model, but the resulting set of equations is somewhat different because of differences in the logical constraints applied: we solve a quadratic system of equations by elimination of intracellular CO_2 concentration (C_i), instead of specifying C_i as an initial condition. Other differences include a more detailed dependence of the kinetic parameters on temperature (Woodrow and Berry 1988), and a simplifying assumption that empirically relates the maximum rate of electron transport to maximum carboxylation velocity (Wullschleger 1993).

Photosynthesis is regulated by canopy conductance to CO₂, leaf maintenance respiration and daily meteorological conditions including air pressure, air temperature and photosynthetically active photon flux density (PPFD). Canopy conductance to CO₂ is calculated as a proportion (62.5%) of canopy conductance to water vapor (g_c) . Maximum canopy conductance to water vapor $(g_{c,max})$ defines the upper boundary of the photosynthetic rate and is determined by LAI and prescribed leafscale boundary layer, cuticular and maximum stomatal conductances to water vapor; gc is reduced when air temperature, vapor pressure deficit, PPFD, or soil water potential deviates from prescribed optimal conditions (Jarvis and Morison 1981, Running and Coughlan 1988, Leuning 1990). BI-OME-BGC represents the canopy as a "big leaf" in that all units of leaf area in the canopy are represented by a single, canopy-averaged conductance. This assumption is generally

not valid at sub-daily (e.g., hourly) time-steps because the reduction in irradiance at lower vertical layers of the canopy reduces conductances at the bottom of the canopy. The big leaf assumption is strengthened, however, by the integrative effects of a daily time-step, and by the implicit assumption that allocation of leaf nitrogen between light harvesting and carbon fixing enzymes over depth in the canopy varies in response to the canopy light environment, allowing an optimized use of intercepted radiation (Evans 1989).

Total respiration from the system (R_{tot}) is estimated on a daily basis as the sum of the maintenance (R_m) , heterotrophic (R_h) and growth (R_g) respiration components. Maintenance respiration represents the total loss of carbon due to day and night leaf respiration $(R_{dl} + R_{nl})$, sapwood (R_{sw}) , coarse root (R_{cr}) and fine root (R_{fr}) respiration. Respiration is estimated as a daily proportion of carbon in living tissue that is released as the result of cellular metabolic processes, excluding any growth processes. Maintenance respiration is calculated from mean daily air temperatures and prescribed leaf, root and sapwood carbon pools, using an exponentially increasing function of respiration with temperature as described by Amthor (1986). The magnitude of respiration response to temperature is governed by a prescribed rate defined at a reference temperature (i.e., 20 °C) and a proportional change in rate for a 10 °C change in temperature (Q_{10}) . In all cases except leaf maintenance respiration, the daily average temperature is used in calculating a single value for the mass lost to maintenance respiration for the day. However, for leaves, $R_{\rm dl}$ and $R_{\rm nl}$ rates are calculated from estimated day and night air temperatures, respectively, because $R_{\rm dl}$ is required to determine GPP. Daily growth respiration (R_g) was not determined explicitly by the model, but was computed as a proportion (32%) of the daily difference between GPP and $R_{\rm m}$ (Penning de Vries et al. 1974).

The heterotrophic respiration term in BIOME–BGC represents the loss of carbon from the system resulting from soil microbial respiration. Daily R_h is estimated as a proportion of prescribed soil and litter carbon pools. The proportion of litter carbon being respired on a daily basis is regulated by soil water potential and soil temperature, as described by Orchard and Cook (1983), Andren and Paustian (1987), and Running and Coughlan (1988). The proportion of soil carbon respired on a daily basis was estimated as 1% of the proportion of litter carbon respired (Cole and Rapp 1981, Fox and Van Cleve 1983).

Net primary production (NPP) represents the net accumulation of carbon by the stand and is determined as the difference between GPP and the sum of $R_{\rm m}$ and $R_{\rm g}$. The net ecosystem exchange of carbon (NEE) represents the net accumulation or loss of carbon by the entire soil–stand system and is determined as the difference between GPP and $R_{\rm tot}$. In this investigation, positive fluxes denote a net uptake of carbon by the system and negative fluxes denote a net carbon loss.

BIOME–BGC uses daily maximum and minimum air temperatures, humidity, incident solar radiation and precipitation to determine daily carbon and water fluxes. Average daily incident shortwave radiation (R_s) was simulated by MT–CLIM

(mountain microclimate simulator) logic as described by Running et al. (1987). Average daily net solar radiation ($R_{\rm n,s}$) was estimated from a prescribed, constant albedo for vegetation. Average daily net solar radiation was attenuated through the vegetation canopy using Beer's formulation and a prescribed extinction coefficient modulated by LAI to derive the amount of solar radiation transmitted through the canopy ($R_{\rm s,t}$). The amount of solar radiation absorbed by the canopy ($R_{\rm s,a}$) was estimated as the difference between $R_{\rm n,s}$ and $R_{\rm s,t}$. The PPFD was estimated based on the assumption that photosynthetically active radiation represents approximately 50% of $R_{\rm s,a}$ (Running and Coughlan 1988).

Mean daily air temperature (T_a) was estimated as the average of the measured daily maximum and minimum air temperatures. Minimum daily air temperature was assumed equal to mean daily dew point and was used with T_a to estimate mean daily vapor pressure deficit (VPD). Mean daily soil temperatures within the prescribed effective soil depth ($T_{\rm soil}$) were estimated based on an 11-day running average of T_a (Zheng et al. 1993). Soil water potential (PSI) was estimated from soil water content, soil depth and texture information described by Cosby et al. (1984). The variables T_a , VPD, PPFD and PSI were used to estimate g_c and GPP, as described by Jarvis and Morison (1981) and Farquhar and von Caemmerer (1982), respectively. Both T_a and $T_{\rm soil}$ were used to estimate $R_{\rm m}$, and $T_{\rm soil}$ and PSI were used to estimate $R_{\rm m}$, and $T_{\rm soil}$ and PSI were used to estimate $R_{\rm m}$, and $T_{\rm soil}$ and PSI were used to estimate $R_{\rm m}$, and $T_{\rm soil}$ and PSI were used to estimate $R_{\rm m}$, and $T_{\rm soil}$ and PSI were used to estimate $T_{\rm soil}$ 0 and PSI were used to estimate $T_{\rm soil}$ 1 and PSI were used to estimate $T_{\rm soil}$ 2 and PSI were used to estimate $T_{\rm soil}$ 3 and PSI were used to estimate $T_{\rm soil}$ 4 and $T_{\rm soil}$ 5 and PSI were used to estimate $T_{\rm soil}$ 6 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 3 and PSI were used to estimate $T_{\rm soil}$ 4 and $T_{\rm soil}$ 5 and PSI were used to estimate $T_{\rm soil}$ 5 and PSI were used to estimate $T_{\rm soil}$ 6 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate $T_{\rm soil}$ 8 and PSI were used to estimate

Study sites

The BOREAS study region consists of a $1 \times 10^6 \text{ km}^2$ area covering portions of central Saskatchewan and Manitoba (Sellers et al. 1995). Within this region are two intensive study areas, each approximately 10,000 km² in area. The Southern Study Area (SSA) is located near Prince Albert, Saskatchewan (53.2° N, 105.7° W), whereas the Northern Study Area (NSA) is located roughly 500 km to the northeast near Thompson, Manitoba (55.7° N, 97.8° W). The BOREAS study region consists of relatively flat to gently rolling terrain with mean elevations from 261 m in the NSA to 520 m in the SSA. Vegetation cover is predominantly coniferous with low species diversity and productivity relative to temperate forests. Understory vegetation is composed of sparse shrubs with extensive moss and lichen ground cover. Temperature and solar radiation play a major role in controlling boreal forest productivity by limiting energy, water supply and nutrient availability to plants for much of the year. The growing season is generally restricted to the months between May and September when daily air temperatures rise above 0 °C.

Tower eddy flux and micro-meteorological measurement networks were established within selected forest stands considered representative of the boreal forest region. Detailed descriptions of site characteristics and data collections are presented by Sellers et al. (1995), Black et al. (1996) and Baldocchi et al. (1996). Two tower sites were examined within the NSA representing young jack pine (NYJP) and old black spruce (NOBS) forest types. Three tower sites were examined

within the SSA consisting of old black spruce (SOBS), old jack pine (SOJP) and old aspen (SOAS) forest types.

The jack pine sites were 30–80% forest covered, with jack pine (*Pinus banksiana* Lamb.) ranging in ages from less than 15 years at the NYJP site to between 70 and 90 years at the SOJP site. Canopy heights ranged from 10–15 m at the SOJP site and from 4–5 m at the NYJP site. Understory vegetation was sparse, consisting predominantly of isolated groups of alder (*Alnus crispa* (Ait.) Pursch) with an extensive surface cover of lichens (*Cladina* spp.), bearberry (*Arctostaphylos uva-ursi* (L.) K. Spreng) and bog cranberry (*Vaccinium vitisideae* L.). Soils were coarse textured, sandy and well drained, and were classified as a degraded Eutric Brunisol/Orthic Eutric Brunisol.

The SOBS and NOBS sites were 30–70% forest covered with mature stands of black spruce (*Picea mariana* (Mill.) BSP) from 60–80 years in age with tree heights ranging from 7–18 m. Soils consisted of poorly drained silt and clay predominantly covered by a thick moss layer (*Sphagnum* spp., *Pleurozium* spp.).

The SOAS site was 30–60% forest covered with an extensive stand of mature aspen (*Populus tremuloides* Michx.), approximately 70 years old, with average tree heights of approximately 21 m. The SOAS also had an extensive understory composed primarily of hazelnut (*Corylus cornuta* Marsh.), approximately 2 m in height, interspersed with alder and shrubs. Soils consisted of Orthic Gray Luvisol with a medium to moderately fine silty clay texture.

Measurements

During 1994, air temperature, solar radiation and precipitation were measured at approximately 15-min intervals at each study site (BOREAS Science Team 1995). Detailed descriptions of the methods and instruments used to obtain these data are provided by Sellers et al. (1995), Fan et al. (1995), Baldocchi et al. (1996) and Black et al. (1996). The 1994 meteorological records for each study site were incomplete because of instrument malfunction and other factors. Continuous meteorological records for 1994 were obtained for each study site by temporally interpolating missing data or substituting data from adjacent sites. Daily maximum and minimum air temperatures, precipitation and solar radiation were then derived from the continuous data records for each site and used to generate model results.

Model simulations were compared with estimates of various components of daily and annual carbon budgets derived from CO₂ flux and biomass measurements to determine whether results were consistent. Model simulations of daily NEE were compared with daily carbon fluxes derived from above-canopy CO₂ flux measurements collected over selected periods during 1994 by tower flux teams at each study site (BOREAS Science Team 1995). Canopy CO₂ fluxes were measured by an eddy-covariance flux method. The 24-h integrated CO₂ fluxes were derived from approximate, half-hourly integrated flux measurements and a model for night respiration based on soil and air temperature measurements to adjust for CO₂ storage effects (Baldocchi and Vogel 1996, Black et al. 1996). These data were

obtained from the tower flux teams at each study site and converted to carbon fluxes.

Aboveground NPP estimates (NPP_{ag}) were derived from BIOME–BGC simulations of the 1994 total NPP using above-to belowground NPP ratios reported in the literature for young and mature spruce, pine and aspen stands in boreal and cold temperate environments (Malkonen 1974, Paavilainen 1980, Gower et al. 1992, Wieder and Yavitt 1994). These results were compared with the 1994 NPP_{ag} estimates derived from annual stem-wood radial increment cores and site specific allometric equations for the SOAS, SOBS, SOJP, and NOBS forest types (BOREAS Science Team 1995).

Model initialization

A list of critical parameters used to define soil and stand carbon characteristics at the five study sites is presented in Table 1. These parameters were obtained from both published literature for these genera and unpublished site observations from BOREAS researchers. When data were not available for these genera, parameters were obtained from related genera under similar environmental conditions.

Effective soil depth defines the vegetation rooting depth and is used with soil texture information to determine the maximum amount of soil water available for evapotranspiration. Effective soil depth was set at 0.5 m following Haddeland and Lettenmaier (1995) and Sellers et al. (1995). Soil volumetric water capacity, field capacity and *b*-parameter values were derived from measurements collected at the sites during 1994 by Cuenca et al. (1997) and values reported in the literature for representative soil types (Cosby et al. 1984).

Mean daily vegetation solar albedos for snow-free conditions were estimated from site observations reported by Sellers et al. (1995), whereas the mean daily solar albedo for snow was set at 0.8 (Brutsaert 1988). Canopy extinction of solar radiation was set at -0.5 based on data for temperate deciduous and evergreen forests (Jarvis and Leverenz 1984, Hunt and Running 1992).

Mean annual leaf area index values were derived from overstory LAI measurements conducted over approximately three periods during the 1994 growing season at each study site by Chen (1996). No attempt was made to configure the model to represent understory canopy contributions to daily carbon exchange or seasonal changes in LAI. Specific leaf area (SLA) and leaf nitrogen concentrations were estimated from plucked

Table 1. The BIOME-BGC physiological and soil parameters for the BOREAS study sites.

Parameter	NOBS	NYJP	SOBS	SOJP	SOAS
Leaf area index (one-sided)	4.51	1.68	3.84	2.40	2.70
Specific leaf area (m ² kg ⁻¹ C)	6.0	6.0	6.0	6.0	22.5
Leaf carbon (kg m ⁻²)	0.75	0.28	0.64	0.40	0.12
Sapwood carbon (kg m ⁻²)	0.28	0.06	0.12	0.15	0.40
Coarse root carbon (kg m ⁻²)	0.07	0.02	0.03	0.04	0.10
Fine root carbon (kg m ⁻²)	1.50	0.56	1.28	0.80	0.18
Litter carbon (kg m ⁻²)	0.62	0.53	0.56	0.62	1.00
Soil carbon (kg m ⁻²)	8.7	4.0	8.7	1.4	2.9
Leaf nitrogen (kg m ⁻²)	0.004	0.003	0.004	0.004	0.004
Fraction of leaf N in Rubisco (%)	6.0	6.0	6.0	6.0	14.0
Leaf maint. resp. proportion at 20 °C (1/day)	0.002	0.002	0.002	0.002	0.009
$R_{\rm fr}$ proportion at 20 °C (1/day)	0.002	0.002	0.002	0.002	0.009
$R_{\rm sw}$ proportion at 20 °C (1/day)	0.001	0.001	0.001	0.001	0.004
$R_{\rm cr}$ proportion at 20 °C (1/day)	0.001	0.001	0.001	0.001	0.004
Site shortwave albedo (snow-free conditions)	0.1	0.1	0.1	0.1	0.2
Canopy extinction of PAR	-0.5	-0.5	-0.5	-0.5	0.5
Leaf cuticular conductance (mm s ⁻¹)	0.01	0.01	0.01	0.01	0.05
Max. leaf stomatal conductance (mm s ⁻¹)	1.0	1.0	1.0	1.0	5.0
Max. leaf boundary layer conductance (mm s ⁻¹)	0.8	0.8	0.8	0.8	0.8
Optimal temperature for g_c (°C)	15.0	15.0	15.0	15.0	15.0
Max. temperature for g_c (°C)	40.0	40.0	40.0	40.0	40.0
PSI at start of g_c reduction (MPa)	-0.5	-0.5	-0.5	-0.5	0.5
PSI at complete g_c reduction (MPa)	-1.7	-2.0	-1.7	-1.7	-2.3
VPD at start of g_c reduction (kPa)	1.0	1.0	1.0	1.0	1.0
VPD at complete g_c reduction (kPa)	4.0	4.0	4.0	4.0	4.0
Q_{10} for $R_{\rm m}$	2.3	2.3	2.3	2.3	2.3
Effective soil depth (m)	0.5	0.5	0.5	0.5	0.5
Volumetric soil water capacity (%)	48	40	48	40	48
Soil field capacity (%)	44	10	40	11	34
Available soil water (mm)	220	50	200	57	170
Soil <i>b</i> -parameter	10.39	2.79	10.39	2.79	10.39

needle and leaf measurements at the black spruce, jack pine and aspen sites (Dang et al. 1997). The amounts of leaf nitrogen in Rubisco were taken from the literature for representative cover types (Field and Mooney 1986, Fan et al. 1995). Leaf carbon was derived from LAI and SLA information. Sapwood carbon was estimated from sapwood biomass measurements collected at the BOREAS study sites (BOREAS Science Team 1995), and estimates of the relative proportions of sapwood live cells (Waring and Schlesinger 1985). Coarse root carbon was estimated to be approximately 25% of sapwood carbon (Grier and Logan 1977, Grier et al. 1981).

The amount of carbon attributed to fine root biomass is highly variable depending on species type, stand age and nutrient availability. Processes governing the partitioning of carbon between root and foliar biomass are generally poorly understood. Observations have shown, however, that fine root biomass is generally greater than foliar biomass in nutrientlimited systems, which often occur in boreal and cold temperate forests, and may represent an adaptation to maximize nutrient uptake (Tetreault et al. 1978, Keyes and Grier 1981, Nadelhoffer et al. 1985). Soil carbon attributed to fine roots was estimated from 1.5 (SOAS) to 3.5 (SOJP) times the estimated foliar carbon based on observations from boreal and cold temperate coniferous and deciduous stands on nutrient poor sites (Linder and Axelsson 1982, Perala and Alban 1982, Comeau and Kimmins 1989, Gower et al. 1992). Soil litter and mineralized organic carbon pools within the prescribed soil depths were estimated from soil layer depth, bulk density and percent organic carbon measurements conducted at each of the study sites (BOREAS Science Team 1995).

Leaf, stem, coarse and fine root maintenance respiration coefficients reflect daily rates of carbon respired per unit of carbon in living tissue and were estimated from measured rates for coniferous and deciduous cover types (Vowinckel et al. 1975, Johnson-Flanagan and Owens 1986, Sprugel et al. 1995). Leaf cuticular, boundary layer and maximum stomatal conductances were obtained from Nobel (1991) and Waring and Schlesinger (1985). These parameters show a wide range of values reported in the literature, varying significantly with life-form, boundary layer thickness, drought and shading conditions. Values used for this investigation represent the lower end of the range reported in the literature, based on evidence that reduced nutrient availability in boreal regions causes reductions in carbon uptake and stomatal conductances (Shulze et al. 1994, Baldocchi and Vogel 1996). Temperature, VPD and PSI bounds on g_c were estimated from values reported in the literature for coniferous and deciduous stands (Vowinckel et al. 1975, Waring and Franklin 1979, Waring and Schlesinger 1985, Baldocchi and Vogel 1996).

The Q_{10} parameter quantifies the slope of the exponential relationship between temperature and maintenance respiration. Values of Q_{10} show considerable variation in the literature (e.g., 1.3–3.0), but center near a value of approximately 2.0 for most plants (Sprugel et al. 1995). A Q_{10} of 2.3 was used for all sites based on observations by Ryan (1991) that the Q_{10} value may be higher for woody plants than for herbaceous plants. A

value of 2.3 is generally consistent with Q_{10} values reported by Baldocchi and Vogel (1996) for the jack pine stand in the SSA.

BIOME–BGC was run over a two-year period at each study site. The model was initialized using AMS mesonet station meteorological data for 1989 from the Thompson airport (55.8° N, 97.9° W) for study sites in the NSA, and Prince Albert airport (53.2° N, 105.7° W) and Waskesiu Lake (53.9° N, 106.1° W) for study sites in the SSA (BOREAS Science Team 1995). All analyses of model results in the second year were made using the 1994 meteorological data.

Results and discussion

Meteorological characteristics

The 1994 daily meteorological data records for the study sites showed that the NSA was slightly colder than the SSA, with a mean annual air temperature of -1.6 °C (SD = 16) compared to 1.0 °C (SD = 15) within the SSA. The growing season was defined as the number of days with minimum daily air temperatures above 0 °C. The growing season occurred over a period of approximately 151 days in the NSA and 178 days in the SSA. Mean daily air temperatures during the growing season were similar between the NSA and SSA, averaging 11.7 °C (SD = 6.2). Mean daily vapor pressure deficits for the same period averaged approximately 0.6 kPa (SD = 0.34) in both regions. Annual precipitation for 1994 averaged 44.7 and 42.1 cm for sites in the NSA and SSA, respectively. The greater amount of annual precipitation in the NSA was the result of greater winter snowfall relative to the SSA; however, the SSA received approximately 36 cm of rainfall during the growing season, which was almost 10 cm greater than the NSA received.

BIOME--BGC simulations

Mean daily plots of BIOME-BGC simulations of the 1994 carbon balance for sites within the SSA and NSA are presented in Figures 1 and 2, respectively. Black spruce and SOJP sites generally had low rates of GPP ranging from 0 to 7.6 and 0 to 5.8 g C m⁻² day⁻¹, respectively, over the year, with mean daily fluxes of 4.2 and 3.2 g C m⁻² day⁻¹, respectively, during the growing season. Total respiration rates at these sites were relatively large ranging from 0.3 to 5.3 g C m⁻² day⁻¹ over the year and averaging 3.4 (black spruce) and 2.8 (SOJP) g C m⁻² day⁻¹ during the growing season. Gross primary production exceeded R_{tot} over a 172-day period from April 9 to September 28 at the SOBS site, whereas the SOJP site was a net carbon sink for a 162-day period. Gross primary production exceeded total respiration at the NOBS site over a 139-day period from May 4 to September 20. This period was generally shorter at the NOBS site than at the SOBS site because of the shorter growing season. Net carbon fluxes at the black spruce and SOJP sites ranged from -2.0 to 1.9 g C m⁻² day⁻¹ between winter and summer months. During the growing season, NEE averaged 0.6 (NOBS), 0.9 (SOBS) and 0.4 (SOJP) g C m⁻² day⁻¹. The near-zero daily NEE fluxes reflect the relatively narrow margin between GPP and R_{tot} .

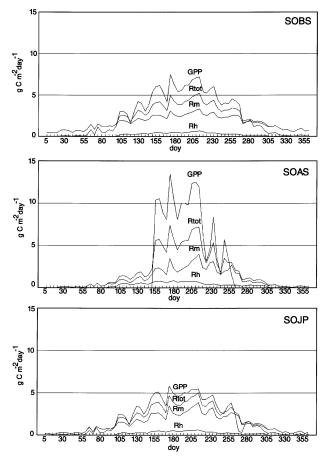


Figure 1. Five-day means of 1994 BIOME–BGC simulated gross primary productivity (GPP), maintenance ($R_{\rm m}$), heterotrophic ($R_{\rm h}$) and total ($R_{\rm tot}$) respiration for the SSA study sites.

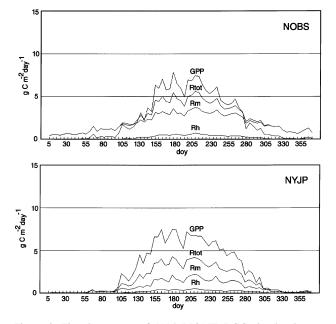


Figure 2. Five-day means of 1994 BIOME–BGC simulated gross primary productivity (GPP), maintenance ($R_{\rm m}$), heterotrophic ($R_{\rm h}$) and total ($R_{\rm tot}$) respiration for the NSA study sites.

Simulated total respiration rates at the SOAS site were relatively large, ranging from 0 to 7.4 g C m $^{-2}$ day $^{-1}$ over the year and averaging 4.8 g C m $^{-2}$ day $^{-1}$ during the growing season. Gross primary production exceeded total respiration at the SOAS site over a 100-day period from May 31 to September 9. The large respiration costs and short period of carbon uptake at the aspen site were generally compensated for by a large photosynthetic capacity, whereby GPP averaged 8.0 g C m $^{-2}$ day $^{-1}$ with maximum daily fluxes of 13.5 g C m $^{-2}$ day $^{-1}$ during the 1994 growing season. The resulting NEE ranged from -2.2 to 6.1 g C m $^{-2}$ day $^{-1}$ over the year and averaged 3.1 g C m $^{-2}$ day $^{-1}$ during the growing season.

Gross primary production at the NYJP site averaged 4.7 (SD = 2.1) g C m⁻² day⁻¹ during the growing season with maximum rates of approximately 7.5 g C m⁻² day⁻¹. Mean daily total respiration rates at this site ranged from 0 to 4.3 g C m⁻² day⁻¹ over the year and averaged 2.6 g C m⁻² day⁻¹ during the growing season. Respiration rates were lower in the NYJP stand than in the older conifer stands because of lower foliar, sapwood and root biomass (Table 1). Gross primary production exceeded total respiration over a 191-day period from April 9 to October 17 and resulted in mean daily NEE ranging from –0.2 to 3.5 g C m⁻² day⁻¹ over the year and averaging 2.0 g C m⁻² day⁻¹ during the growing season.

Heterotrophic respiration simulations represented a small component of $R_{\rm tot}$ at all of the sites, averaging approximately 0.2 (SD = 0.3) g C m⁻² day⁻¹ over the year, whereas $R_{\rm m}$ generally averaged 49 (NYJP) 54 (SOAS), 76 (black spruce) and 77 (SOJP) percent of the total daily respiration rate during the growing season. Maintenance respiration also constituted a large proportion of simulated daily productivity at the SOBS, NOBS and SOJP sites, averaging 73, 88 and 94% of growing season GPP, respectively. The large proportion of mean daily GPP represented by $R_{\rm m}$ at the SOJP site was primarily a result of reduced GPP in response to low soil water content at the end of the growing season. Compared with the SOJP site, $R_{\rm m}$ constituted smaller proportions of daily productivity at the SOAS and NYJP sites, averaging 43 and 30% of GPP, respectively.

On an annual basis, simulated GPP ranged from 7.1 Mg C ha^{-1} year⁻¹ at the SOJP site to 9.6 Mg C ha^{-1} year⁻¹ at the NYJP site (Table 2). Because of lower maintenance respiration costs, estimated total respiration costs only constituted 57% of GPP at the NYJP site, whereas R_{tot} costs at the black spruce, SOAS and SOJP sites were large, representing 79 to 99% of GPP.

For 1994, simulated maintenance respiration costs ranged from 2.7 (NYJP) to 6.8 (NOBS) Mg C ha⁻¹ year⁻¹. Roots constituted more than half of the total maintenance respiration costs, followed by foliage and sapwood $R_{\rm m}$ components. Root respiration also averaged more than twice annual foliar respiration costs. These results concur with annual autotrophic respiration estimates derived from chamber measurements of foliage, woody tissue, and fine roots at BOREAS aspen, black spruce and jack pine sites during 1994 (Ryan and Lavigne 1997).

The large simulated total respiration rates at the black spruce and SOJP sites resulted in a small ($\leq 1.0 \text{ Mg C ha}^{-1} \text{ year}^{-1}$)

estimated net annual carbon accumulation at these sites for 1994. Simulated NEE was greater (1.8 Mg C ha⁻¹ year⁻¹) at the SOAS site than at the black spruce and SOJP sites because of lower annual maintenance respiration costs as a result of the shorter growing season of the deciduous canopy. Among the sites, the NYJP site showed the highest simulated NEE at 4.1 Mg C ha⁻¹ year⁻¹ primarily because of lower maintenance respiration costs.

Model comparison with measured results

BIOME–BGC simulations of daily NEE were compared with NEE derived from stand eddy flux measurements collected during 1994 (Table 3). For all sites, the sample periods for the measured data ranged from 81 to 229 days and averaged 133 (SD = 50.5) days. Sample periods were limited by sensor malfunctions and the high cost and difficulty of maintaining continuous measurements over long periods.

Table 2. The BIOME–BGC predicted annual (1994) carbon budgets (Mg C ha $^{-1}$ year $^{-1}$).

	SOAS	SOBS	SOJP	NYJP	NOBS
GPP ¹	8.7	9.1	7.1	9.6	8.8
$R_{\rm dl}$	0.5	1.1	0.6	0.4	1.2
$R_{\rm nl}$	0.2	0.7	0.3	0.2	0.8
$R_{\rm sw}$	1.2	0.5	0.6	0.2	0.6
$R_{\rm cr}$	0.4	0.1	0.1	0.1	0.2
R_{fr}	1.6	3.6	3.4	1.9	4.0
$R_{\rm m}$	3.9	6.1	5.1	2.7	6.8
$R_{\rm h}$	1.0	0.7	0.7	0.6	0.7
$R_{\rm g}$	2.0	1.3	0.8	2.2	1.1
$R_{\rm tot}$	6.9	8.1	6.6	5.5	8.7
NPP _{tot}	2.8	1.7	1.1	4.7	0.8
NPP_{ag}	2.1	1.3	0.8	2.4	0.6
NEE	1.8	1.0	0.5	4.1	0.1

Abbreviations: GPP = gross primary production; $R_{\rm dl}$ = day foliar respiration; $R_{\rm nl}$ = night foliar respiration; $R_{\rm sw}$ = sapwood respiration; $R_{\rm cr}$ = coarse root respiration; $R_{\rm fr}$ = fine root respiration; $R_{\rm m}$ = maintenance respiration ($R_{\rm dl} + R_{\rm nl} + R_{\rm cr} + R_{\rm fr}$); $R_{\rm h}$ = heterotrophic respiration; $R_{\rm g}$ = growth respiration; $R_{\rm tot}$ = total respiration ($R_{\rm m} + R_{\rm g}$); NPP tot = total net primary production (GPP + $R_{\rm m} + R_{\rm g}$); NPP ag = aboveground NPP; NEE = net ecosystem carbon exchange (GPP - $R_{\rm tot}$).

Table 3. Summary of simulated and tower eddy flux daily carbon fluxes (g C m^{-2} day⁻¹) for 1994 for the BOREAS study sites. RMSE = root mean squared errors.

Site	BIOME-BGC average	Tower average	RMSE	Sample period (Days)
SOJP	0.77	0.55	1.13	81
SOBS	1.60	1.48	1.33	107
SOAS	1.19	1.05	1.59	229
NOBS	0.28	-0.03	0.99	129
NYJP	2.76	1.88	1.69	119

Mean daily differences between simulated and measured NEE ranged from 0.12 to 0.3 g C m $^{-2}$ day $^{-1}$. Root mean squared errors (RMSE) between tower flux measurements and BIOME–BGC results ranged from 0.99 (NOBS) to 1.69 (NYJP) g C m $^{-2}$ day $^{-1}$ and averaged 1.35 g C m $^{-2}$ day $^{-1}$ for all sites. Simulated NEE explained only 51.7% (SE = 1.32 g C m $^{-2}$ day $^{-1}$) of the variance in daily tower flux results for all sites (Figure 3). Simulated results also tended to overestimate tower flux results by approximately 25%.

Differences between BIOME–BGC and tower-flux measured NEE results were probably a result of several factors. One factor is that NEE was simulated on a daily basis as a residual between two values of similar magnitude (Figures 1 and 2). Daily simulations of $R_{\rm tot}$ and GPP were generally within 1.5 g C m⁻² day⁻¹ of each other and were each approximately 3.5 times the magnitude of simulated NEE. These results reflect the age-related decline in productivity observed in older stands (Ryan et al. 1997). The effects of relatively small errors in the estimation of daily GPP or $R_{\rm tot}$ appear to be compounded in daily NEE simulations. For example, simulated changes of \pm 1% in either GPP or $R_{\rm tot}$ resulted in corresponding changes of 143% in mean daily NEE during the growing season at the NOBS site.

Another factor may be the way in which 24-h CO₂ fluxes were determined from tower eddy flux measurements. At night, the CO₂ measurements were problematic because of CO₂ storage effects resulting from low wind conditions at the BOREAS study sites (Baldocchi and Vogel 1996, Black et al. 1996). An empirical function between night CO₂ flux and soil temperature measurements at a 2 cm depth was used with daytime CO₂ flux measurements to calculate 24-h CO₂ fluxes above the SOAS overstory. This function explained 67% of the variance in measured CO₂ fluxes on nights with winds above 3.5 m s⁻¹ (Black et al. 1996). Missing CO₂ flux data were also estimated as an empirical function of photosynthetic photon flux density measurements during the day and air temperature

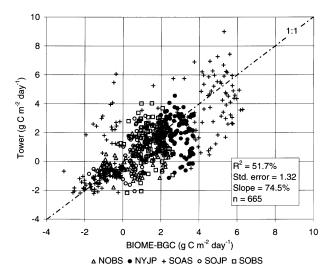


Figure 3. Linear regression relationship between BIOME-BGC and tower flux daily NEE.

at night above the SOJP overstory (Baldocchi and Vogel 1996). In BIOME–BGC, mean daily air temperatures are used to calculate foliar and sapwood respiration, and estimated soil temperatures are used to determine root and heterotrophic respiration; these variables are then used with GPP to calculate NEE. Both BIOME–BGC and tower flux results rely to a large extent on empirical assumptions to determine daily (24-h) NEE at the BOREAS study sites. There is generally considerable scatter in temperature–respiration relationships within forests because of spatial heterogeneity in stand morphology and meteorological conditions, measurement error and possible additional unknown controls on respiration (Wofsy et al. 1993, Fan et al. 1995, Black et al. 1996).

Site heterogeneity may be another factor responsible for differences between measured and model results. BIOME-BGC is a process-level model designed to be general enough to apply at regional to global scales. Several simplifying assumptions are used regarding stand and meteorological conditions to facilitate application at regional scales. A fundamental model assumption is that stand physiological conditions such as age, stand structure, LAI and carbon pools are spatially and temporally uniform on an annual basis. Soil conditions such as depth, density and water content are also assumed spatially uniform, with no lateral or subsurface drainage. Stand conditions at the study sites were both spatially and temporally diverse and were composed of different age types, biomass densities and species compositions (BOREAS Science Team 1995). Leaf area was represented in BIOME-BGC using a constant, seasonal average LAI based on overstory measurements conducted at the beginning, middle and end of the 1994 growing season at each study site. However, LAI measurements showed approximate seasonal variations of 12, 18 and 34% of mean values for black spruce, jack pine and aspen stands, respectively. Additionally, understory contributions to carbon fluxes were not represented in model simulations because only overstory LAI was used to characterize canopy leaf area. Observations by Black et al. (1996), however, indicate that the hazelnut understory had a significant effect on seasonal carbon fluxes at the SOAS site.

BIOME–BGC and tower flux NEE results also reflected different time-scales of meteorological data which may have contributed to poor agreement between daily results. Daily tower flux measurements were derived from integrated half-hourly measurements and reflect physiological responses to sub-daily meteorological conditions; leaf stomata and canopy CO₂ conductances, for example, are capable of responding to environmental conditions at time-scales on the order of minutes. BIOME–BGC, however, is driven by daily averaged radiation, temperature and VPD, and reflects an averaged g_c response.

Five-day means of daily results were compared to evaluate seasonal trends in simulated and measured NEE. Plots of these results are presented for SSA and NSA sites in Figures 4 and 5, respectively. The RMS errors between 5-day means of measured and modeled NEE ranged from 0.58 to 1.49 g C m⁻² day⁻¹. Model results also accounted for 66.1% of the variation

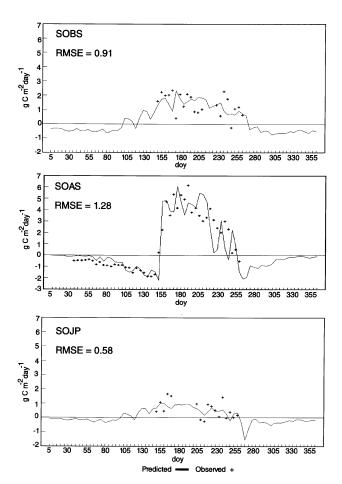


Figure 4. Five-day means of 1994 BIOME-BGC and tower flux NEE for the SSA study sites

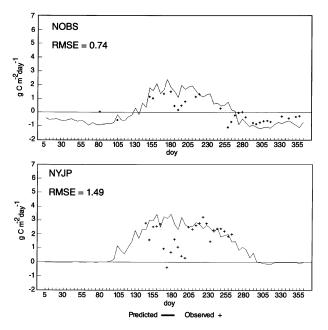


Figure 5. Five-day means of 1994 BIOME-BGC and tower flux NEE for the NSA study sites.

in tower flux NEE (SE = $0.97 \text{ g C m}^{-2} \text{ day}^{-1}$) when 5-day means of daily results were compared.

At the SOAS site, model results showed large reductions in NEE during the period from approximately August 7 to September 3 in response to simulated soil water stress which the tower-flux measurements did not detect. This period generally coincided with the observed onset of aspen and hazelnut canopy senescence (Black et al. 1996). Simulated soil water contents during this period ranged as low as 25% and were within the range of measured conditions (Cuenca et al. 1997, Kimball et al. 1997). Tower eddy flux measurements above the aspen canopy showed maximum daily evapotranspiration rates from 3.5 to 4.0 mm day⁻¹ during the 1994 growing season (Black et al. 1996); BIOME-BGC simulations of daily evapotranspiration were generally consistent with 1994 tower measurements at the SOAS site (Kimball et al. 1997). Tower measurements and simulations both showed that evapotranspiration rates above the aspen canopy were generally larger than at the BOREAS conifer tower sites during the 1994 growing season. This enhanced depletion of soil water in SOAS simulations led to water stress near the end of the growing season and a reduction in NEE. The magnitudes of the reductions in BIOME-BGC and tower NEE differed and may reflect a greater volume of water available to the aspen stand than simulations show, possibly through lateral transport of water from adjacent soil layers or deeper rooting depths. Simulations using effective soil depths up to 1 m (i.e., maximum amount of available soil water = 340 mm), however, still showed droughtrelated reductions in NEE near the end of the 1994 growing season.

At the SOJP site, BIOME–BGC NEE simulations also showed reductions over two periods in response to water stress. These periods occurred from approximately August 23 to September 5 and from September 13 to October 3. Simulated soil water contents during these periods averaged approximately 4% and were also within the ranges of measured conditions (Cuenca et al. 1997, Kimball et al. 1997). Tower flux NEE data were not available at the end of September for comparison. However, Baldocchi and Vogel (1996) concluded that low soil water contents severely limited evapotranspiration rates from the SOJP stand and it is likely that CO₂ fluxes were similarly affected because of the strong linkage between photosynthesis and stomatal conductance (Collatz et al. 1991).

At the NYJP site, differences between model and tower NEE results were particularly large from approximately June 13 to July 19 (Figure 5). The weather during this period was generally cool and wet. BIOME–BGC simulations showed reductions in both GPP and $R_{\rm tot}$ in response to reduced photosynthetic rates and lower air temperatures (Figure 2). The net effect on NEE was therefore minimal. Tower NEE values during this period may have been more sensitive to the soil temperature data that were used to estimate CO₂ fluxes at night. Simulated daily soil temperatures were generally warmer than air temperatures during this period. Warmer soil temperatures may have resulted in enhanced night respiration rates and reduced tower-measured NEE values relative to BI-OME–BGC results.

Measured aboveground, overstory NPP at the tower sites and the ranges of NPP $_{\rm ag}$ at adjacent sites with similar stand characteristics were approximately 2.1 \pm 1.01, 1.3 \pm 0.52, 0.8 \pm 0.24 and 0.6 \pm 0.62 Mg C ha $^{-1}$ year $^{-1}$ at the SOAS, SOBS, SOJP and NOBS sites, respectively, based on tree increment core data and allometric equations for 1994 (BOREAS Science Team 1995). Differences between measured and simulated NPP $_{\rm ag}$ ranged from 0.01 to 0.7 Mg C ha $^{-1}$ year $^{-1}$. No NPP $_{\rm ag}$ estimates were made from biomass measurements at the NYJP site, though comparison with measurements made at a mature, high productivity jack pine site within the SSA was approximately 1.3 Mg C ha $^{-1}$ year $^{-1}$ for 1994, yielding a difference of 1.0 Mg C ha $^{-1}$ year $^{-1}$ with model results.

Conclusions

Comparisons between BIOME-BGC simulations and field data were encouraging given the general nature of model parameterizations of stand morphology and our current, limited understanding of boreal forest structure and processes. Model simulations accounted for approximately half of the variance in observed daily carbon fluxes; differences between tower and simulated NEE were attributed to several factors including problems measuring nighttime CO2 fluxes and model assumptions of site homogeneity. Comparisons between simulations and field data improved markedly, however, at coarser time-scales. Simulated NEE accounted for 66.1% of the variance in tower measurements when 5-day means of daily fluxes were compared. Simulations of annual NPP ag were also within the ranges of results derived from tree increment core and allometric equations, and annual simulations of the components of $R_{\rm m}$ were in general agreement with results determined from chamber measurements of foliage, sapwood and roots. Replication of field measurements and confidence in simulations are likely to improve as we learn more about stand structure and the various processes controlling water and carbon fluxes within the boreal forest from BOREAS field campaigns.

Simulated results showed that all of the sites were small net carbon sinks for 1994. Younger stands were more productive than older stands, though most of these sites appear to be less productive than temperate forests (Grier and Logan 1977, Kinerson et al. 1977, Edwards et al. 1981, Waring and Schlesinger 1985). Our results also imply that these boreal forest stands currently do not have the capacities to be the large terrestrial carbon sinks that global atmospheric CO₂ measurements and general circulation models indicate should occur in temperate and boreal latitudes of the Northern Hemisphere (Tans et al. 1990, Ciais et al. 1995). More work is needed, however, to determine annual carbon budgets within other boreal stands, regional patterns of carbon exchange and processes that regulate carbon fluxes from stand to regional scales.

Stands with large daily evapotranspiration rates and stands situated on sandy soils with low water holding capacities may be particularly vulnerable to increased temperature and drought conditions. Site productivity simulations at the SOAS site were strongly limited by low soil water conditions near the

end of the growing season because of enhanced water loss through large daily evapotranspiration rates. Simulated SOJP site productivity was also adversely affected by low soil water conditions because of coarse, sandy soils at the site, though the SOJP canopy appears better able to conserve water through lower daily evaporative water loss. Stands subject to frequent water stress during the growing season may exist on the margin between being annual net sources or sinks for atmospheric carbon, especially older stands that exhibit low photosynthetic and high respiration rates.

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